

# High Resolution Simulation of Nonstationary Random Fields

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# Stochastic Weather Generators

- ▶ Agricultural, ecological, hydrologic models often require daily weather (e.g. precipitation, minimum/maximum temperature, solar radiation)
  - ▶ On grid
  - ▶ In the future
- ▶ **Stochastic Weather Generators (SWGs)** can be used to produce infinitely long series of synthetic weather, for observation network infilling, or climate model downscaling
- ▶ SWGs are statistical models whose simulated values “look like” observed weather
  - ▶ Daily statistics
  - ▶ Interannual statistics
  - ▶ **Spatial** statistics
- ▶ Approaches: **empirical**, **model-based**, (combination?)

# Spatial Models

Modern SWGs require **spatial** simulations that honor observed spatial correlations.

Difficulties:

- ▶ **Simulation over complex terrain**
- ▶ Simulation over large heterogeneous areas

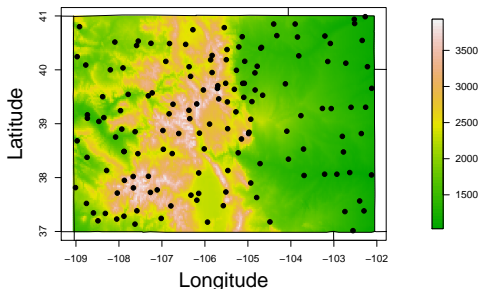
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Data: 145 stations from the Global Historical Climatology Network. Daily minimum temperature during 1893-2011.



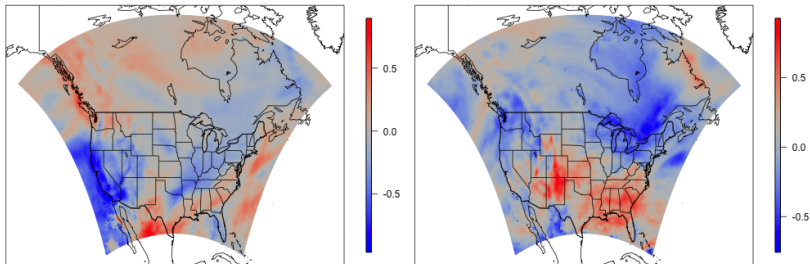
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More data: Gridded average winter precipitation anomalies from a regional climate model output (11,760 grid cells).



# Properties of $Z(\mathbf{s})$

If  $Z(\mathbf{s}), \mathbf{s} \in \mathbb{R}^d$  is a **Gaussian** process, then we need only specify

- ▶ The **mean** function  $\mathbb{E}Z(\mathbf{s})$  (suppose  $= 0$ )
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Usual approach: Simulate independent normals  $\varepsilon$  and form  $L\varepsilon$ , where  $L$  is a Cholesky factor.

**Main problem:**  $Z(\mathbf{s})$  is (often) a nonstationary process, computing  $L$  is **difficult**.

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Suppose  $Z(\mathbf{s})$  is a nonstationary process that can be written

$$Z(\mathbf{s}) = D(\varphi(\mathbf{s}))$$

where

- ▶  $D(\mathbf{s})$  is a stationary process with covariance  $C_D(\cdot)$
- ▶  $\varphi : \mathbb{R}^d \rightarrow \mathbb{R}^d$  is an invertible deformation function.

Then,

$$\begin{aligned}\text{Cov}(Z(\mathbf{s}_1), Z(\mathbf{s}_2)) &= C(\mathbf{s}_1, \mathbf{s}_2) \\ &= C_D(\varphi(\mathbf{s}_1) - \varphi(\mathbf{s}_2)).\end{aligned}$$

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**Key idea:** To simulate a vector  $(Z(\mathbf{s}_1), \dots, Z(\mathbf{s}_n))'$ , simulate  $(D(\varphi(\mathbf{s}_1)), \dots, D(\varphi(\mathbf{s}_n)))'$  via stationary simulation algorithm.

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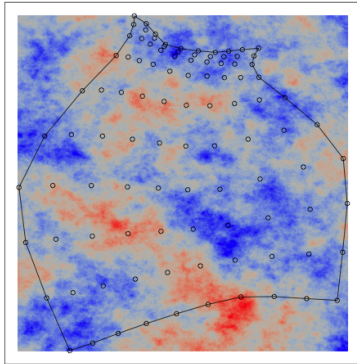
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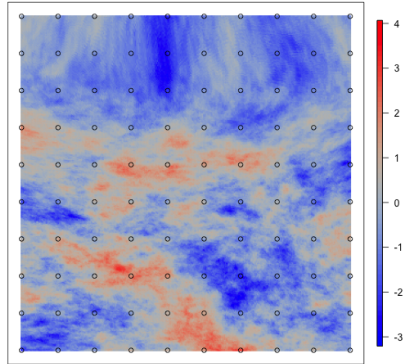
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# Example

Stationary plane



Nonstationary plane



# Stationary simulation: spectral method

The **spectral method** (Shinozuka and Jan, 1972) relies on the representation

$$Z(\mathbf{s}) = \text{Re} \left[ \int \exp(2\pi i \boldsymbol{\omega}' \mathbf{s}) dY(\boldsymbol{\omega}) \right].$$

Simulations follow from a discretization of the integral.

This method is **approximate**.



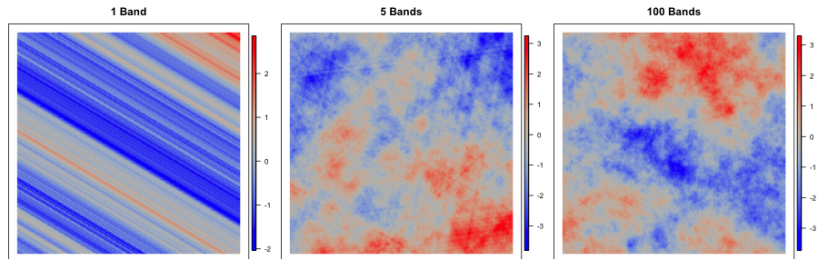
# Stationary simulation: turning bands

The **turning bands** (Matheron, 1973; Mantoglou and Wilson 1982) method uses

$$Z(\mathbf{s}) = \frac{1}{\sqrt{L}} \sum_{i=1}^L Z_i(\mathbf{s} \cdot \mathbf{e}_i)$$

where  $\{Z_i(\cdot)\}_{i=1}^L$  are mutually independent one-dimensional processes,  $\{\mathbf{e}_i\}_{i=1}^L$  are unit vectors.

This method is **approximate**.



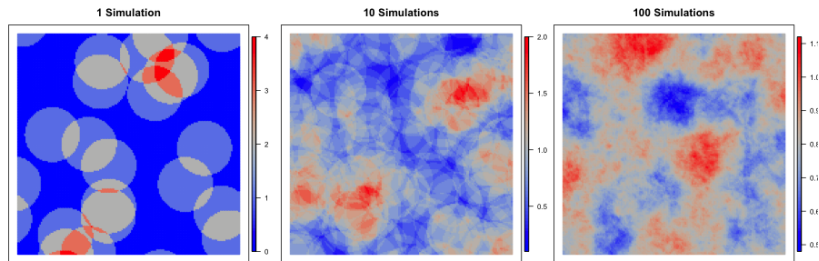
# Stationary simulation: random coin

The **random coin** approach (Chilés and Delfiner 1999; Schlather 2012) requires

$$C(\mathbf{h}) = \int g(\mathbf{s})g(\mathbf{s} + \mathbf{h})d\mathbf{s}.$$

The approximate simulation is  $\sum_{\mathbf{h} \in \Pi} g(\mathbf{s} - \mathbf{h})$  where  $\Pi$  is a stationary Poisson point process with unit intensity.

This method is **approximate**.



# Stationary simulation: Circulant embedding

The **circulant embedding** method (Dietrich and Newsam 1993, 1997; Wood and Chan 1994) requires simulation locations to be on a regular grid.

In  $d = 2$  dimension, if  $\{\mathbf{s}_i\}_{i=1}^n$  are regular then

$$\Sigma_0 = \left( C(\mathbf{s}_i - \mathbf{s}_j) \right)_{i,j=1}^n$$

is block Toeplitz. We **embed**  $\Sigma_0$  in a circulant matrix

$$\Sigma = \begin{pmatrix} \Sigma_0 & \Sigma_{12} \\ \Sigma_{21} & \Sigma_{22} \end{pmatrix} = FDF^*$$

where  $D$  is diagonal with eigenvalues,  $F$  has complex exponential entries.

This method is **exact**.

# Estimation

Need to estimate:

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For observations  $Y(\mathbf{s}_i)$  at locations  $\mathbf{s}_1, \dots, \mathbf{s}_n$ , use

$$\hat{C}(\mathbf{x}, \mathbf{y}) = \frac{\sum_{i=1}^n \sum_{j=1}^n K_\lambda(\|\mathbf{x} - \mathbf{s}_i\|) K_\lambda(\|\mathbf{y} - \mathbf{s}_j\|) Y(\mathbf{s}_i) Y(\mathbf{s}_j)}{\sum_{i=1}^n \sum_{j=1}^n K_\lambda(\|\mathbf{x} - \mathbf{s}_i\|) K_\lambda(\|\mathbf{y} - \mathbf{s}_j\|)}$$

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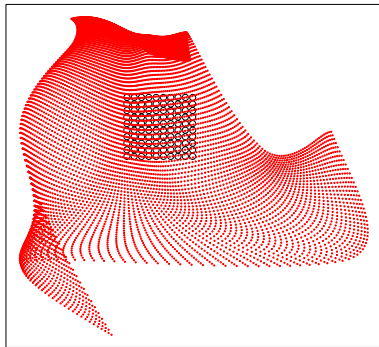
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Estimate  $\varphi$  via

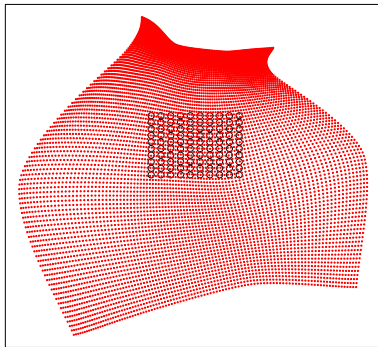
$$\min_{\varphi} \sum_{i=1}^n \sum_{j=1}^n (\hat{C}(\mathbf{s}_i, \mathbf{s}_j) - C_D(\varphi(\mathbf{s}_i) - \varphi(\mathbf{s}_j)))^2 + \lambda \langle \varphi, \varphi \rangle.$$

# Effect of penalty term

**Without Penalty**

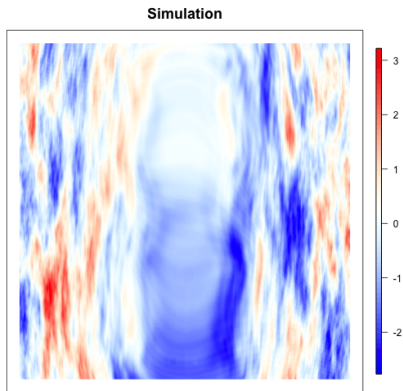
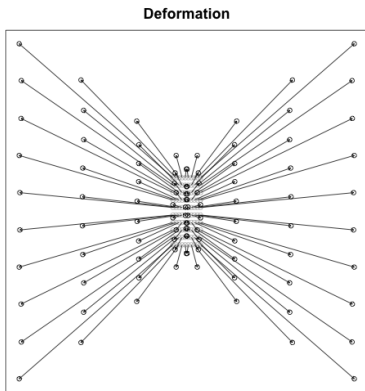


**With Penalty**



# Valleys

A major difficulty is that weather processes often exhibit differing spatial structure in valleys than over higher terrain. Additionally, can allow for continuous simulation between regimes, e.g., land and ocean.

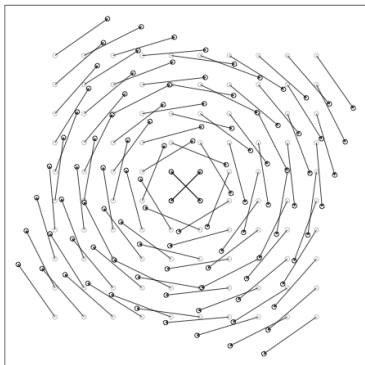




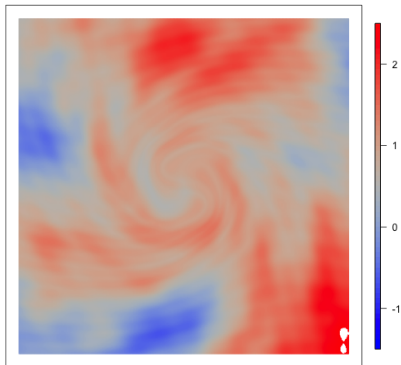
# Vortices

Vortex-like behavior is crucial at microscales for wind modeling as well as macroscales such as for hurricanes.

Deformation



Simulation



Model for observations of minimum temperature  $Y(\mathbf{s})$ ,  $\mathbf{s} \in \mathbb{R}^2$ ,

$$\begin{aligned} Y(\mathbf{s}) &= \mu(\mathbf{s}) + Z(\mathbf{s}) + \varepsilon(\mathbf{s}) \\ &= \text{“Climate”} + \text{“Weather”} + \text{“Error”} \end{aligned}$$

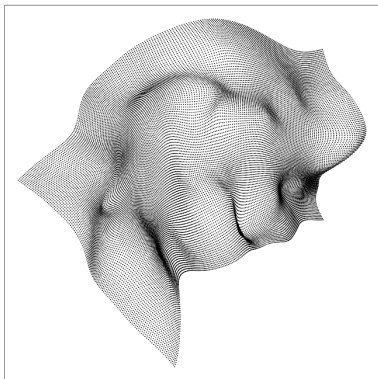
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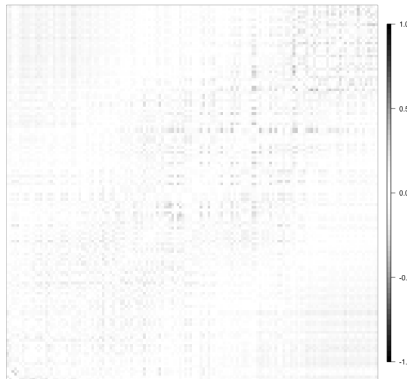
Steps:

- ▶ Extract  $\hat{\mu}(\mathbf{s})$  via linear regression with spatially varying coefficients
- ▶ Estimate covariance  $C(\cdot, \cdot)$  of  $Z(\mathbf{s})$  nonparametrically
- ▶ Suppose the stationary process  $D(\mathbf{s})$  has Matérn covariance with smoothness 1, range estimated from data
- ▶ Fit  $\hat{\varphi}$
- ▶ Simulate  $Z(\mathbf{s}) = D(\varphi(\mathbf{s}))$  on a grid of 25,000 locations

Deformation

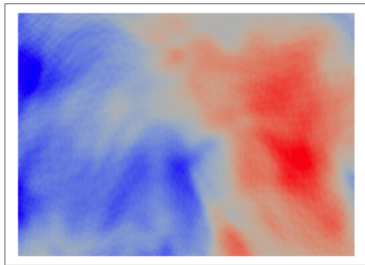


Deformed correlation error

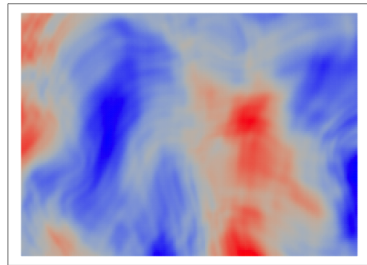


# Colorado

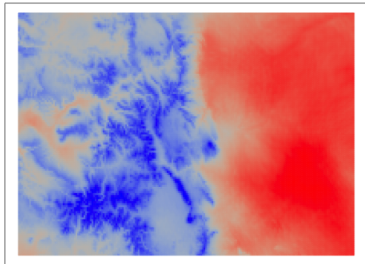
Residual 1



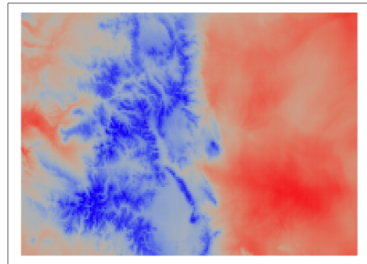
Residual 2



Simulation 1

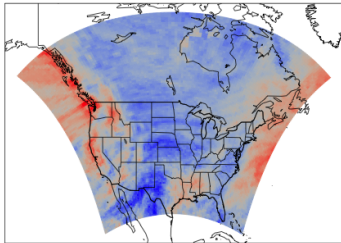


Simulation 2

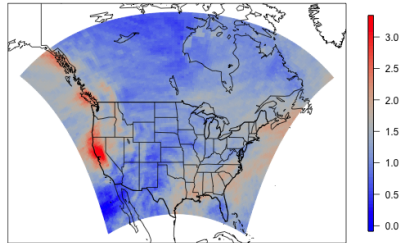


# Regional Climate Model

Simulation 1



Simulation 2



- ▶ High resolution nonstationary simulation via deformation
- ▶ Extensions: space-time, multivariate, non-Gaussian fields (precipitation)
- ▶ Better deformation functions?

# Shameless Plug

Verdin, A., Rajagopalan, B., Kleiber, W. and Katz, R. W. (2014)  
“Coupled stochastic weather generation using spatial and  
generalized linear models.” *Stochastic Environmental Research  
and Risk Assessment*, in press.

- ▶ Simulation of minimum, maximum temperature and precipitation occurrence
- ▶ Temperature coupled to precipitation occurrence in conditional fashion

